1	Dynamics of bacterial communities in relation to soil aggregate
2	formation during the decomposition of ¹³ C-labelled rice straw
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26 Abstract

27 The addition of fresh organic matter is known to modify both microbial community structure and soil aggregation. The objective of this study was to understand the relationship 28 29 between the dynamics of the soil microbial community structure in relation to that of their 30 habitats during the decomposition of straw. Soil samples, ground (< 200 µm) to remove macroaggregates, were amended with uniformly ¹³Clabelled powdered rice straw ($< 500 \mu m$) 31 32 and incubated for 21 days. Unamended control samples were also incubated under the same 33 conditions. Total C and rice straw C (C_{Straw}) mineralised or remaining in different soil fractions 34 $(0-50, 50-200, 200-2000 \text{ and } > 2000 \mu \text{m})$ were measured. Fatty acid methyl ester (FAME) 35 profiling was used to determine total bacterial community structure and FAME based stable isotope probing (FAME-SIP) was used to characterise the straw degrader communities. The 36 mineralisation rate of the native C and the C_{Straw} was high. The formation of macroaggregates 37 38 $(> 2000 \,\mu\text{m})$ occurred within 2 days in amended and unamended samples but did so to a greater 39 extent in the amended samples. The C_{Straw} was mainly located in fractions > 200 μ m, where degraders were the most abundant. The ¹³C-FAME profiles followed the same trends as total 40 41 FAME profiles through time and within soil fractions, suggesting common dynamics between straw degraders and total bacterial communities: Gram-negative were more important in 42 43 fraction $> 200 \ \mu m$ and during the early stages of the incubation while Gram-positive and 44 actinobacteria dominated in fine fractions and at the end of the incubation. Bacterial community 45 structure changed rapidly (within 2 days) in conjunction with the formation of new microbial 46 habitats, suggesting that the relationship between the two is very close.

47

48 **Keywords:** Soil Bacterial Communities Structure, ¹³C-labelling, FAME-SIP, Soil Aggregation,

49 Mineralisation, Microscale Biogeography

50 **1. Introduction**

51 The soil environment is made up of a huge diversity of microenvironments in which the 52 biological component of soil exists and is active (Young and Ritz, 1998). The physical 53 architecture of solid and pore space results in a complex distribution of oxygen, water films and 54 gradients of solutes spanning distances as small as a few micrometers. These 55 microenvironments form the microbial habitats in soil. It is known that the different 56 microenvironments select for different microbial communities. Community structure is 57 dependent on the size of soil aggregates and on the location within soil aggregates (Ranjard et 58 al., 2000; Poly et al., 2001; Mummey and Stahl, 2004), as are a number of physiological 59 characteristics of the communities (Hattori, 1988). Aggregate size classes differ in their 60 potential denitrification (Lensi et al., 1995; Sev et al., 2008), mineralisation of organic matter 61 (Franzluebbers and Arshad, 1997; Sey et al., 2008) and enzymatic activities (Drazkiewicz, 1995). Microbial communities and activities are also known to vary with pore size class (Strong 62 63 et al., 2004; Ruamps et al., 2011). However, the microbial habitat in soil exists in a highly 64 dynamic state as microbial activity (Dorioz et al., 1993; Feeney et al., 2006), plant root growth 65 or alterations in the water status can all affect the physicochemical properties of the 66 microenvironments. Changes in the microbial habitat can also occur due to soil management 67 practices such as tillage (Six et al., 2004; Young et al., 2006). Soils consist of mineral and 68 organic materials bound together to form soil aggregates. Soil aggregates are defined by their sizes and their stability in water (Tisdall and Oades, 1982). Microaggregates (<200 µm) and 69 70 macroaggregates (> 200 µm) regulate key factors such gas and liquid diffusion (Sexstone et al., 71 1985; Horn and Smucker, 2005) and soil organic matter (SOM) turnover (Puget et al., 2000; 72 Six et al., 2004). These changes are all likely to affect the abundance, the structure, and therefore the functioning, of the resident microbial communities (Young and Crawford, 2004). 73

74 The process of aggregation occurs firstly by the formation of macroaggregates (> 20075 µm), which are formed by mineral associations with particulate organic matter (POM) via temporary binding agents (i.e. fungal hyphae and plant roots). Microaggregates (< 200 µm) are 76 77 formed within macroaggregates around POM (Oades, 1984; Six et al., 2000). The 78 decomposition of POM within macroaggregates by microorganisms produces 79 exopolysaccharides and other metabolites which act as persistent binding agents for the 80 formation of microaggregates (as reviewed by Degens, 1997). Thus, the dynamics of soil 81 aggregates create different ecological niches. There is a particularly strong niche differentiation 82 during the formation of macroaggregates around POM, as there are large amounts of readily 83 available substrate. Regions of the soil surrounding POM (extending no further than a few millimeters) have been termed the 'detritusphere' and are described as microbial activity 84 hotspots (Gaillard et al., 1999; Poll et al., 2006). Macroaggregates contain more C, and more 85 86 labile young C, than microaggregates (Puget et al., 2000) and the C turnover is also higher in 87 macroaggregates (Buyanovsky et al., 1994; Monreal et al., 1997). In contrast, microaggregates 88 are more stable and persistent than macroaggregates (Jastrow, 1996; Bossuyt et al., 2002) and 89 are characterised by low predation rates, low nutrient availability and low gas diffusion rates as 90 reviewed by Ranjard and Richaume (2001).

91 The transient nature of soil aggregates makes it legitimate to ask how the dynamics of 92 microbial communities are related to the evolution of their habitat but also how this affects the 93 active communities involved in the degradation of POM. As microorganisms are active agents 94 of aggregate formation and stabilization (Tisdall and Oades, 1982), they impact on their own 95 habitat. The *de novo* formation of macroaggregates provides the best conditions for studying 96 this, as habitat differentiation is likely to be the most significant. Several studies (e.g. Denef et 97 al., 2001, 2002) have investigated the influence of nutrients addition on macroaggregates 98 formation/stabilization on soil in which macroaggregates were crushed (< 250 μ m) prior to

99 incubation. However, none of these studies assessed the response of the microbial communities 100 to macroaggregate formation after OM addition. Accordingly, the objectives of this study, were: 101 (i) to determine whether microbial community structure was related to habitat differentiation 102 during soil aggregate formation (ii) to assess the influence of fresh organic matter addition on 103 microbial dynamics during aggregation and (iii) to identify the location (macroaggregates vs 104 microaggregates) and the structure of the bacterial communities that assimilate C from fresh 105 organic matter into their biomass. To this end, soil without macroaggregates was amended with 106 uniformly ¹³C-labelled rice straw and incubated for 21 days. The ¹³C-labelling was used in order 107 to differentiate native from added residue C and to target bacterial communities using straw C 108 among the total soil bacterial communities using FAME-SIP.

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110 2. Materials and Methods

111 2.1. Soil samples and plant residues

112 The experimental site was located in the Anjozorobe district (18°46'S, 47°32'E), near 113 Antananarivo, Madagascar. The soil studied was a Ferrasol under a vegetated fallow dominated 114 by Aristida sp. The soil texture was constituted by 12% of fine sand (50-200 µm), 28% of 115 coarse sand (> 200μ m), 30% of silt and 30% of clay. The soil was constituted of different soil 116 fractions: $0.2\% > 2000 \mu m$, $84.1\% 200-2000 \mu m$, $13.5\% 50-200 \mu m$ and $2.3\% 0-50 \mu m$. The total organic carbon (TOC) was 23.0 ± 0.6 g kg⁻¹, the total nitrogen was 1.7 ± 0.1 g kg⁻¹ and 117 118 the pH (H₂O) was 5.8. The bulk isotopic signature of the organic carbon was $-12.3 \pm 1.76\%$. 119 In March 2007, samples were collected to a depth of 10 cm from several locations at the site 120 and bulked to form a composite sample of about 3 kg. The soil was air-dried after sampling. 121 Particulate organic matter was removed from the soil with tweezers to reduce the amount of 122 native POM available to microorganisms. About 48% of total POM was removed (0.0047 g POM g^{-1} soil remained). In order to destroy the macroaggregates (> 200 μ m), the soil was 123

124 crushed and sieved (< 200 μ m). The sand that had been removed by the sieving (> 200 μ m) was 125 added to the sieved soil to ensure that the soil texture was not altered.

126 Uniformly ¹³C-labelled rice straw was obtained from the Groupement de Recherches Appliquées en Phytotechnologies (CEA/DSV/DEVM, Cadarache, France). The rice was grown 127 128 in culture chambers in which the atmospheric total CO₂ concentration was kept constant at 360 129 ppm. The CO₂ partial pressure in the chamber was continuously monitored by Near Infrared 130 Spectroscopy. Regulation was achieved by automatic injection of 13 C-enriched CO₂ (10.5 ± 131 0.4%), which started 20 days after sowing. Rice plants were harvested 129 days after sowing. 132 The straw used was a mix of stems and leaves from rice plants, with C and N contents of 417.5 \pm 1.7 g kg⁻¹ and 19.7 \pm 0.3 g kg⁻¹, respectively. The bulk isotopic signature of the rice straw C 133 134 was $6124 \pm 158\%$. Before incorporation into the soil samples, the straw residues were 135 ovendried at 60 °C and finely ground to $< 500 \,\mu$ m.

136 2.2. Experimental design

137 Amended samples (20 g dry weight equivalent) were mixed with 0.08 g of rice straw 138 residues, corresponding to a C addition of 7.3% of the total soil organic C. Control samples did 139 not receive any amendment. Prior to commencing the incubation, the water potential of all 140 samples was adjusted to -0.01 MPa by addition of Milli-Q water. The samples were placed in 141 1.2 l jars along with a vial of Milli-O water (20 ml), to avoid drying of soil, and a vial of NaOH 142 (20 ml, 0.5 M) for trapping the CO₂ released during soil respiration (see Section 2.3). 143 Microcosms were incubated in the dark for 21 days at 30 °C. It was decided to incubate for 21 144 days based on the previous studies showing that microbial communities were mainly affected 145 by the addition of plant residues during the first 15 days (McMahon et al., 2005) or 21 days (Schutter and Dick, 2001) of incubation which lasted 80 days. Six replicates of control and 146 147 amended samples were destructively sampled before the incubation started (day 0), after 2 (day 148 2) and 21 days (day 21) of incubation. Three replicates of control and amended samples were 149 immediately fractionated and three other replicates (unfractionated soil samples) were stored at 150 -20 °C for further analyses.

151 2.3. Soil respiration

The CO₂ produced by soil respiration was determined by back titration at equivalent point pH 8.6 using HCl (0.2 M), of the NaOH placed in the microcosms with an excess of BaCl₂ (1.5 M). After each titration, the carbonate precipitate (BaCO₃) was filtered (fiber glass filter, Wathman GF/A ϕ 5.5 cm, retention 0.26 mm), washed, and dried at 40 °C. The ¹³C content of the BaCO₃ was determined using an EA-IRMS (NA-1500, Carlo-Erba). The amount of ¹³C mineralised was obtained as explained below.

158 2.4. Soil fractionation and analyses

159 The soil fractionation method used was adapted from Yoder (1936). The samples were 160 immersed in Milli-Q water (4 °C) above a sieve (mesh size: 2000 µm) for 5 min. Large 161 macroaggregates (> 2000 µm) were then separated by automatically moving the sieve up and 162 down 3 cm at 30 cycles per minute for 10 min. Aggregates retained on the 2000 µm sieve were 163 considered to be large macroaggregates. The water and soil that passed through the 2000 µm 164 sieve were then sieved using a 200 µm followed by a 50 µm mesh sieve to obtain the 200-2000 165 μm and 50-200 μm soil fractions. The water and soil that passed through the 50 μm sieve was 166 centrifuged (2000 rpm for 10 min, 4 °C) to obtain 0-50 µm fractions. Particulate organic matter 167 (POM) fractions were discarded by flotation in water from the 200-2000 and 50-200 µm soil 168 fractions in order to reduce the amount of plant residues and plant FAME. Each soil fraction 169 was subsampled (from 0.5 g to 2 g) to measure the weight distribution of the soil fractions after oven drying at 40 °C. Subsequently, the subsamples were ground (< 200 μ m) for C and ¹³C 170

- analyses, using an Elemental Analyser (EA, NA-2000, Carlo-Erba) and an Elemental Analyser
- 172 coupled to an Isotopic Ratio monitoring Mass Spectrometer (EA-IRMS, NA-1500, Carlo-Erba),
- 173 respectively. The remaining soil fraction material was stored at -20 °C for FAME analysis.

174 2.5. FAME analyses

FAME profiles were produced following the protocol described by Schutter and Dick, 175 176 (2000). Briefly, fatty acids were extracted and methylated from 3 g of soil (dry weight 177 equivalent) with 15 ml of a mixture of NaOH (0.2 M) and methanol (1 M). The mixture was 178 then vortexed 20 s and shaken for 1 h at 37 °C. After this incubation, 2-3 ml of acid acetic (1 179 M) was added to neutralise the pH. Hexane (15 ml) was added to each tube, vortexed for 1 min 180 and centrifuged (1600 rpm) at 4 °C for 1 h to isolate the FAME in the organic phase. The 181 organic phase (~ 5 ml) was transferred to a clean tube and evaporated under a stream of N₂. Dry 182 FAME at bottom of the tubes were dissolved in 170 µl MTBE (Methyl Tertiary Butyl Ether)-183 hexane (1:1, v:v) and transferred to a vial with 30 µl of methyl stearate (internal standard, 0.01 184 M). FAME were quantified on a Gas Chromatograph (HP 6890) coupled to a Flame Ionisation 185 Detector (GC-FID) and identified by GC (HP 6890) coupled to an Agilent 5973 electronic 186 impact (70eV) quadripole mass spectrometer (GC-MS). The isotopic signature of each FAME 187 was determined using a GC (HP 5890) coupled to an Isochrom III isotopic mass spectrometer 188 (Micromass-GVI Optima) via a combustion interface (GC-c-IRMS). There were three 189 analytical replicates for all of the FAME isotopic analyses. All GCs were equipped with the 190 same SGE BPX-5 column (50m x 0.25mm x 0.32µm). The injector and detector were 191 maintained at 280°C. Column temperature was programmed at 50 °C for 1 min and then ramped at 2 °C min⁻¹ to 350 °C, followed by an isothermal period of 10 min. Samples were injected in 192 193 splitless mode.

194 The fatty acid nomenclature used was that described by Frostegard et al. (1993). Mono-195 unsaturated and cyclopropyl fatty acids were taken as gram-negative bacteria biomarkers 196 (Zelles, 1999), iso- and anteiso-fatty acids as gram-positive bacteria biomarkers (O'Leary and 197 Wilkinson, 1988; Zelles, 1999) and carboxylic acids with a methyl function on the carbon chain 198 as biomarkers for actinobacteria and in particular Nocardia (Zelles et al., 1994). A number of 199 FAME (C12:0, C14:0, C15:0, C16:0, C17:0, C16:4(3,7,11,15), C18:2(9,12), C18:1(9)cis, 200 C18:0, C19:0, C20:0 or C22:0) were detected in the lipid extract of the rice straw with an overall 201 δ^{13} C value of 8209±247‰ (data not shown). These were not used when analysing microbial 202 community structure as their origin (plant or microbial) was uncertain. Thus, 12 FAME that 203 were solely of bacterial origin, were selected: i-C15:0, a-C15:0, i-C16:0, brC17:0, i-C17:0, 204 aC17:0 for gram-positive bacteria; C16:1(9)cis, C16:1(11)cis, cycC17:0, cycC19:0 for gram-205 negative bacteria; 10me-C16:0 and 10meC18:0 for actinobacteria.

206 2.6. Estimating the amount of rice derived C within each soil fraction

207 The standard notation for expressing isotopic contents as δ and the international internal 208 standard Vienna PeeDee Belemnite ($R_{VPDB} = 0.0112372$) were used (Coplen, 1995). The 209 methylation of isolated fatty acids introduced one additional carbon which was not present in 210 the parent compound and which altered the isotopic ratio of the original carboxylic acids. The 211 measured isotopic ratios of the FAME were corrected for the isotope ratio of the methyl 212 esterified moiety in order to obtain the isotope ratios of the fatty acids as described by Lerch et 213 al. (2007). In the amended soils (AS), equations (1), (2) and (3) were used to estimate the 214 quantity of C_{Straw} mineralised from bulk soil and the amount of C_{Straw} incorporated into each soil 215 fraction and into the FAME.

 $216 \qquad C_{AS} = C_{SOM} + C_{Straw}$ (1)

217
$$\delta^{13}C_{AS} \times C_{AS} = \delta^{13}C_{SOM} \times C_{SOM} + \delta^{13}C_{Straw} \times C_{Straw} (2)$$

where C_{AS} is the total amount of C in a given fraction of the amended soil, C_{SOM} is the amount of C derived from the soil organic matter (SOM) and C_{Straw} is the amount of C derived from the rice straw, $\delta^{13}C_{AS}$ is the isotopic signature of the amended soil, $\delta^{13}C_{Straw}$ is the isotopic signature of rice straw (6124‰) and $\delta^{13}C_{SOM}$ is the isotopic signature of the initial SOM. Here, the $\delta^{13}C$ of control samples was used to estimate $\delta^{13}C_{SOM}$. The amount of C_{Straw} incorporated into a given fraction (i.e. CO_2 evolved, bulk soil, physical fraction or FAME) was determined with equation (3), which is a combination of equations (1) and (2):

$$C_{Straw} = C_{AS} \times \left(\frac{\delta^{13} C_{AS} - \delta^{13} C_{CS}}{\delta^{13} C_{Straw} - \delta^{13} C_{CS}} \right)$$
(3)

225

226 2.7. Statistical analyses

Differences in bacterial community structure were analysed by principal components analysis. Significant differences among samples or among the scores of the PCA were tested by ANOVA (using the day of sampling and the physical fraction as factors). The normality and homoscedasticity of data were checked prior to statistical analysis. The evenness (*E*) of the labelled and total FAME profiles were compared to determine how the incorporation of labelled C was spread throughout the microbial communities. The evenness of FAME profiles was calculated as:

$$E = \left(-\sum_{i=1}^{q} p_i \log p_i\right) / \log q \tag{4}$$

where *p* is the relative abundance of each FAME (molar percentage) and *q* is the total number
of FAME (Legendre and Legendre, 1998). All statistical analyses were performed using R
version 2.9.0 (R Development Core Team. 2009. R: a language and environment for statistical
computing).

3. Results

240

3.1. Dynamics of soil aggregation and carbon distribution among aggregates

241 During the incubation there was a shift towards macroaggregates, with a significant (P 242 < 0.001) decrease in the quantity of fractions $< 2000 \mu m$ and the *de novo* formation of 243 macroaggregates (Fig. 1). The process occurred more rapidly in the amended soil but by day 244 21 the aggregate size distributions of control and amended soil were not significantly different 245 (P > 0.05; Fig. 1). The carbon distribution profiles did not follow the same pattern (Fig. 2). First of all, there were no significant differences between control and amended samples, with the 246 exception of a slight, but significant (P < 0.05), difference in the 200–2000 µm fraction, at 247 sampling day 2. There was a significant (P < 0.01) decrease in the C concentration of the 248 249 fractions < 200 µm during the incubation, with a concomitant increase in the C concentration 250 of the fractions > 200 μ m. However, at all sampling dates, the C concentration in the 0–50 μ m fraction was higher than that of the other fractions. Initially, the C from the straw (C_{Straw}) was 251 252 present mainly in the fractions $< 200 \ \mu m$ (Fig. 3). The C_{Straw} concentration of the fractions >253 200 µm increased dramatically by day 2. In contrast, the C_{Straw} concentration of the other 254 fractions decreased. After 21 days of incubation, all the soil fractions showed a decrease in 255 C_{Straw} , with the exception of the 0–50 μ m fraction in which the C_{Straw} concentration remained 256 steady. It should be noted that the amount of C_{Straw} incorporated into each soil fraction would be 24% lower when using the average δ^{13} C signature of plant fatty acids (8209‰) rather than 257 258 the bulk (6125‰).

260

^{259 3.2.} Carbon mineralisation

The total soil respiration was always significantly higher in amended soil than in control 262 263 soil (data not shown). The difference in the total CO₂ between the non-amended soil and the amended soil was explained by the mineralisation of 13 C-labelled rice residues (79 ± 7 and 793 264 \pm 19 µg ¹³C–CO₂ g⁻¹ soil, at days 2 and 21, respectively). In the amended soil, approximately 265 half of the initial amount of rice straw was mineralised after 21 days of incubation (Table 1). 266 The cumulative mineralisation of the soil organic matter in non-amended soil was always 267 significantly (P < 0.05) higher than in amended soil. The rate of C-SOM mineralisation for non-268 269 amended and amended soil was $0.37 \pm 0.01\%$ and $0.33 \pm 0.01\%$ of C-SOM per day, 270 respectively, after 2 days of incubation and then decreased to $0.24 \pm 0.01\%$ and $0.21 \pm 0.01\%$ of C-SOM per day, respectively. The mineralisation rate of C_{Straw} was more constant than rate 271 272 of C-SOM through the incubation with $2.4 \pm 0.22\%$ and $2.3 \pm 0.06\%$ of C_{Straw}, after 2 and 21 273 days of incubation, respectively.

274 3.3. Dynamics of FAME profiles among soil aggregates

Fatty acid methyl ester profiles showed clear differences among soil fractions. These 275 276 differences were evident principally along the first ordination axis of the PCA (Fig. 4). Lipid 277 profiles of soil fractions 200–2000 and $> 2000 \mu m$ were characterised by a high relative 278 abundance of FAME characteristic of Gram-negative bacteria, such as C16:1(11)cis, 279 C16:1(9)cis and cycC19:0, whereas the 50-200 µm and the 0-50 µm soil fractions were 280 characterised by relatively more branched saturated FAME, characteristic of Gram-positive 281 bacteria, such as aC15:0, iC15:0, aC17:0, iC16:0, i-C17:0 10me-C16:0, brC17:0 and 10me-282 C18:0. The second ordination axis revealed temporal changes in bacterial community structure 283 (Fig. 4). In all soil fractions there was a trend towards a higher relative abundance of branched 284 FAME on day 21 of the incubation, although ANOVA showed that this trend was only 285 significant for the soil fractions $> 50 \,\mu\text{m}$. There was no significant effect of amendment except for the larger (> 200 μ m) fractions after 2 days. 286

Molecular isotope analyses of selected FAME extracted at day 0 showed similar ¹³C 287 288 signature between control and straw amended soils, confirming that the groups of FAME 289 analysed were not from plant origin. At days 2 and 21, all the selected FAME in the amended samples were significantly ¹³C enriched (between 23.1‰ and 313.2‰) relative to the control 290 291 soils (between -13.6‰ and -12.7‰), indicating an incorporation of C from labelled straw into 292 the microbial biomass. The structure of bacterial communities that incorporated C_{Straw} into their 293 biomass differed significantly (P < 0.001) among soil fractions (Fig. 5). Communities in 294 macroaggregates differed from communities in microaggregates and there were further 295 differences between communities associated with 0-50 and 50-200 µm fractions. Inspection of 296 the PC loadings indicated that the lipids responsible for the differences among aggregate size 297 fractions were the same for total and labelled lipids (Fig. 5B): Gram-negative bacteria in 298 macroaggregates (> 200 μ m) and Gram-positive bacteria in microaggregates (< 200 μ m). Significant changes (P < 0.05) in community structure were observed between days 2 and 21 299 300 in all soil fractions except for $0-50 \mu m$, although the changes were not consistent across 301 aggregate-size fractions. Fig. 6 shows that the evenness of the control and amended sample 302 lipid profiles were similar, with the exception of the 0–50 fraction and fractions > 200 μ m at 303 day 2. The FAME profiles were always more even for the total soil (0.98 ± 0.1) than for each 304 soil fraction. While the total and the $< 200 \,\mu m$ fractions did not vary with time, the largest soil 305 fractions had a similar evolution, with a lower evenness value after 2 days. Fig. 7 shows that the evenness of ¹³C-labelled FAME profiles were always lower than those of total FAME. In 306 307 contrast to the total soil, changes in evenness occurred for all the soil fractions during the 308 incubation. With exception of the $0-50 \mu m$ fraction, the labelled FAME profiles were more 309 even after 21 days than after 2 days of incubation. The FAME-C_{Straw} to C_{Straw} ratio was used as 310 an indicator of the amount of biomass relative to the available resource in the habitat. The ratio 311 of FAME-C_{Straw} to C_{Straw} (Fig. S1) was always significantly lower in the finest fraction (0-50

312 μ m) than in the other fractions, with the exception of the > 2000 μ m fraction ratio, which 313 decreased after 21 days of incubation.

314 **4. Discussion**

315 4.1. Carbon Mineralisation

316 Nearly half of the added C_{Straw} was mineralised after only 21 days. This rate is higher 317 than those reported in other studies using similar ¹³C-labelled plant residues. For instance, 318 Bernard et al. (2007) found that 33% of wheat straw residues were mineralised after 28 days 319 incubation. The difference is unlikely to be due to N availability as the C/N ratio was low 320 compared to other studies (Nicolardot et al., 2001), and no mineral N was added. The 321 experimental conditions might explain the high mineralisation rates as the mineralisation of 322 native C was also very high (5% in 21 days). The most likely reason is that both the soil and 323 the straw residues were air dried and crushed prior to incubation, which probably reduced the 324 physical protection of the organic matter (OM), resulting in enhanced mineralisation of both 325 native and added C (Chevallier et al., 2004; Salome et al., 2010). Furthermore, the soil was 326 rehydrated before incubation, which may have caused the desorption of organic compounds, 327 making them available for mineralisation (Van Gestel et al., 1991; Chevallier et al., 2004; Lerch 328 et al., 2011). Chevallier et al. (2004) found 2.4 times more C was mineralised after 35 days of 329 incubation for a soil that was air-dried and rewetted prior to incubation than for a fresh soil. The high temperature of incubation used in this study (30 °C) may also explain the high rate of 330 331 degradation (Dalias et al., 2001).

332 It should be noted that the quantification of the C_{Straw} mineralisation may have been 333 over-estimated because of the ¹³C-labelling method used. Indeed, it is difficult to ensure a 334 totally homogenous ¹³C-labelling of the entire plant because of the isotopic fractionation that 335 naturally occurs during plant metabolism (Ghashghaie et al., 2003). This hypothesis was confirmed when measuring the δ^{13} C values of fatty acids extracted from the rice straw (more than 8209‰) compared to the overall bulk value (around 6125‰). If the more ¹³C enriched molecules were also the most labile, the amount of C_{Straw} would have been overestimated. This potential bias might also explain why an apparent negative priming effect (less native C mineralised in the amended soil compared to the control soil, Table 1) was observed during the incubation.

342 *4.2. Formation of de novo soil aggregates*

343 The incubation of crushed soil resulted in the rapid formation of *de novo* soil aggregates 344 in both treatments after only 2 days. Soil macroaggregate formation may have been induced by 345 a number of factors in both treatments. The wetting of dry soil, OM adsorption to clays, the 346 physical entanglement of soil particles by fungi and the production of polysaccharide binding 347 agents during microbial decomposition of OM are all known to cause aggregation (Tisdall and 348 Oades, 1982; Bossuyt et al., 2001). Aggregate formation likely also occurred because of 349 physical or electrostatic interactions between the clay minerals and oxides (Oades and Waters, 350 1991; Denef et al., 2004). Oxides are the main binding agent in Ferralsols (Oades and Waters, 351 1991). Despite the major role of oxides in the Ferralsol aggregation, the addition of plant 352 residues increased the rate of aggregation with time (Fig. 1). The absence of significant 353 difference between amended and non-amended soil aggregation after 21 days can be explained 354 by the presence of oxides in both treatments and, to a lesser extent, the presence of native 355 organic debris in the non-amended soil (only 48% were removed) which might also indirectly 356 promote soil macroaggregation. In contrast, the amount of smaller soil fractions (< 200 μ m) 357 decreased with time, suggesting that the latter fractions were incorporated into macroaggregates 358 as described in the aggregate hierarchy model (Oades and Waters, 1991; Bossuyt et al., 2001; 359 Denef et al., 2001). Such de novo rapidly formed macroaggregates were likely microhabitats

with specific physico-chemical characteristics, such as organic resources, water and oxygenavailability.

362 4.3. Dynamics of total bacterial communities

363 The analysis of microbial community structure using FAME profiling to distinguish 364 microbial communities from different soils is now well established and gives results that are 365 similar to DNA-based methods (Ritchie et al., 2000; Schutter and Dick, 2000). Although the 366 extraction of total fatty acids can include non-microbial fatty acids, the method used in this 367 study, the ester-linked (EL) procedure, reduces the risk of this happening (Schutter and Dick, 368 2000). Schutter and Dick (2002) showed that plant amendments did not elevate eukaryote fatty 369 acids above background levels 3 days after amendment. This was also the case here. No 370 difference in the fungal biomarker C18:2(9,12) was observed after the addition of rice straw 371 and the significantly higher C18:2(9,12) content observed in fractions $> 200 \,\mu\text{m}$ of the amended 372 soil after 21 days of incubation (data not shown) can be attributed to an enhanced development 373 of fungi in larger aggregates after the addition of straw. It is now firmly established that fungi 374 participate in aggregate formation through hyphal growth (e.g. Oades and Waters, 1991; 375 Degens, 1997; Bossuyt et al., 2001) and this result corroborates what has been found before. 376 However, the isotopic signature of this fungal biomarker (> 800‰, data not shown) suggested 377 that there was a contribution of rice straw lipids to the FAME profiles measured, despite the 378 fact that this was not reflected in total lipid contents. Moreover, neutral and glycolipid fatty 379 acids can end up in Eukaryote peaks when lipids are analysed by the EL-FAME method 380 (personal communication R.P. Dick). If there was a high production rate of neutral and 381 glycolipids during the decomposition of the rice straw, these lipids may have masked the fungal 382 marker signal. The ambiguous origin of certain lipids such as C18:2(9,12) meant that ecological 383 interpretation of the profiles was not straightforward. Therefore, the potential impact of non-384 microbial fatty acids being included in the analyses was eliminated by removing all fatty acids

detected in the lipid extract of the rice straw. Therefore, the differences among fractions and as
a function of time in both total FAME and ¹³C-FAME profiles (Figs. 5 and 6) were likely caused
by changes in bacterial communities rather than by plant derived FAMEs.

388 The rapidity with which the bacterial community structure in the *de novo* formed 389 macroaggregates diverged from the community structure of the other fractions was totally 390 unexpected. Biological communities are known to be structured through interactions with their 391 habitat (O'Malley and Dupre, 2007). This is particularly true for microbial communities: 392 microbial biogeographic studies have shown that contemporary environmental effects (i.e. 393 habitat) play a significant role in structuring microbial communities (Martiny et al., 2006; King 394 et al., 2010). Here, there was a clear differentiation in bacterial community structure among soil 395 fractions, confirming the results of others (Fall et al., 2004; Mummey and Stahl, 2004) and 396 suggesting that there is a microscale biogeography in soil (Mummey et al., 2006; Ruamps et 397 al., 2011). The differences in bacterial community structure were primarily due to higher 398 relative abundances of actinobacteria and other Gram-positive bacteria in the fractions < 200 399 µm and greater proportions of Gram-negative bacteria in macroaggregates, again in agreement 400 with previous studies (e.g. Mummey and Stahl, 2004). The higher relative abundance of Gram-401 negative bacteria in the present study might result from their role in aggregation, through the 402 production of extracellular polysaccharides: Amellal et al. (1998) showed that water-stable 403 aggregates of rhizosphere soil could be increased by inoculating the soil with Gram-negative 404 bacteria. The differentiation in bacterial community structure might also be due to differences 405 in local conditions in the soil fractions and, in particular, the uneven distribution of organic 406 matter between soil aggregates (Elliott, 1986; Puget et al., 1995). Gram-negative bacteria are 407 thought to be r-strategist/copiotrophic organisms with a high growth rate, using labile substrate; 408 actinobacteria and most Gram-positive bacteria on the other hand, are thought to be K-409 strategist/oligotrophic organisms that are better decomposers of less labile soil organic matter

410 but have a lower growth rate (Fierer et al., 2007; Elfstrand et al., 2008). More C derived from 411 the rice straw was found in the macroaggregates, indicating that there was a more readily 412 available C source for r-strategists such as the Gram-negative bacteria mainly found within 413 macroaggregates. Finally, sample preparation may also have had an effect on the rapid change 414 in bacterial community structure in the *de novo* formed macroaggregates. The soil was crushed 415 $(< 200 \text{ }\mu\text{m})$ to remove all the macroaggregates, possibly resulting in microbial cell death and 416 the release of labile organic matter (Balesdent et al., 2000). If the labile organic matter thus 417 released was primarily available to microorganisms in the *de novo* formed macroaggregates, 418 this might explain the development of rapidly growing r-strategist type bacteria in these 419 aggregates. Bacterial community succession was similar in all aggregate size fractions, with a 420 higher relative abundance of Gram-positive and actinobacteria at the end of the incubation. This 421 pattern of microbial succession, where rapidly growing r-strategist organisms are replaced by 422 slower growing K-strategist organisms n the latter stages of organic matter decomposition, has 423 been previously described (Fierer et al., 2007; Bastian et al., 2009; Baumann et al., 2009; Lerch 424 et al., 2011). This study also produced the first evidence for greater variation in bacterial 425 community structure over time in large aggregate size fractions (> 200 μ m). These fractions 426 were also the only ones in which the total C concentration (Fig. 2) and the straw-C concentration 427 (Fig. 3) varied during the incubation, suggesting that the amplitude of change in bacterial 428 community structure is closely related to changes in the microbial habitat

429 4.4. Dynamics of straw degrader communities

All the FAMEs studied were ¹³C-labelled after only 2 days for the amended soil, indicating that the straw was rapidly incorporated into the bacterial biomass. This is consistent with the rapid mineralisation observed during the same period and has been observed by others previously (McMahon et al., 2005). The distribution of ¹³C-labelling in FAME showed that both Gram-negative and Gram-positive bacteria were involved in the decomposition, which 435 again confirmed the results of McMahon et al. (2005). Using molecular methods for taxonomic identification of bacteria (DNA- and RNASIP), Bernard et al. (2007) showed that the 436 assimilation of ¹³C from wheat residues was composed exclusively of Gram-negative bacteria 437 after 14 days incubation. The difference in results may be due to the different methodological 438 439 approaches, soil and plant residue differences aside. DNA- or RNA-SIP required a very high degree of ¹³C-labelling compared to lipid-SIP in order to separate "heavy" from "light" nucleic 440 441 acids by density centrifugation and microorganisms that incorporated low levels of ¹³C may 442 have remained undetected by DNA-SIP. The fact that Gram-negative bacterial lipids presented between 2 and 19 times more ¹³C-labelling than lipids characteristic of Gram-positive bacteria 443 444 in the $> 50 \,\mu\text{m}$ soil fractions and unfractionated soil corroborates the assertion that DNA-SIP may not be sensitive enough to detect low levels of labelling. 445

446 Principal component analysis showed that the degrader community structure followed 447 the same general trends as the whole unlabelled community: relatively more Gram-negative 448 bacteria in the largest fractions and an increase in the relative abundance of Gram-positive 449 (including actinobacteria) in all size fractions during the incubation (Fig. 5). Although all the FAMEs analysed were enriched in ¹³C, the evenness of the ¹³C enriched lipid profiles relative 450 451 to the total lipid profiles (Figs. 6 and 7) suggested that the specificity of the bacterial degraders 452 was higher than the total bacterial community. These results are in agreement with that obtained 453 by Bernard et al. (2007). The lower evenness found in the largest size fractions (> 200 µm) also 454 suggested a higher specificity of degraders compared to the whole biomass. The evenness of 455 the labelling increased in the largest fractions between 2 and 21 days (Fig. 7), probably due to 456 the occurrence of cross-feeding (Gallagher et al., 2005) and also simply because nearly all soil 457 was present in macroaggregates by day 21. Furthermore, the FAME-C_{Straw}/C_{Straw} ratio (Fig. S1), 458 used as an indicator of the amount of biomass relative to the available resource in the habitat, 459 decreased for the largest soil fractions between 2 and 21 days of incubation. The fraction 200460 2000 μ m showed the lowest evenness values after 21 days of incubation for the total and the 461 degrader communities. The low evenness of the community in this fraction seems to be constant 462 throughout the incubation, which might highlight a more stable habitat, in term of nutrient 463 availability, than the > 2000 μ m fraction.

464 **5.** Conclusions

This study provides a basis for understanding how the dynamics of microbial habitat 465 466 formation through soil aggregation and fresh organic matter mineralisation drives microbial 467 community structure. The study clearly shows that the variation in bacterial community 468 structure was greater in the *de novo* formed habitats (large aggregate size fractions > $200 \,\mu$ m) 469 than in the pre-existing ones (size fractions $< 200 \ \mu$ m). Furthermore, the *de novo* formed 470 habitats tended to have a greater proportion of fast growing organisms (i.e. r-strategists). These 471 results are consistent with Odum's theory of ecosystem succession (Odum, 1969), which 472 suggests that mature system tend to be more stable than the young ones and that young systems 473 are characterised by high growth rates. Differentiation of bacterial community structure among 474 different habitats was demonstrated to occur very rapidly (i.e. 2 days), indicating that bacterial 475 community structure is closely linked to habitat. Using ¹³C-labelling substrates, it was 476 demonstrated that the active component of the bacterial communities, the straw degrader 477 communities, were mainly present in, but not limited to, the macroaggregates.

478 **References**

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662 Figures and tables

Table 1: Cumulative amounts of C_{SOM} mineralised in control and amended soil expressed in percentage of initial C_{SOM} , and cumulative amount of C_{Straw} mineralised in amended soil expressed in percentage of initial C_{Straw} . Standard deviations were calculated for 3 experimental replicates.

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670	Source of C mineralised	2 days	21days
671	%C _{SOM} in control soil	0.75±0.03	5.0±0.13
672 673	%C _{SOM} in amended soil	0.67±0.02	4.5±0.21
075	% C _{Straw} in amended soil	4.80±0.44	48.2±1.17
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Aggregate Distribution (g 100 g⁻¹ soil)

Fig. 1. Weight distribution of the soil fractions, in control soil (grey bars) and amended soil (dark grey bars) through time. Error bars represent the standard deviation of 3 experimental replicates. * show a significant difference between control and amended soil for a specific date and soil fraction (P < 0.05).



Fig. 2. Total organic carbon concentration of the soil fractions, in control soil (grey bars) and amended soil (dark grey bars) through time. Error bars represent the standard deviation of 3 experimental replicates. * show a significant difference between control amended soil for a specific date and soil fraction (P < 0.05).





Fig. 3. Carbon concentration derived from the rice straw added (C_{straw}) in the soil fractions of
the amended soil through time. Error bars represent the standard deviation of 3 experimental
replicates.

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Fig. 4. Scores (A) and loadings (B) of the 2 main components of the PCA, representing 81% of the variability. Variables are the relative abundance of 12 FAME detected between Day 0 (white), Day 2 (grey) and Day 21 (black) of the incubation for each soil fractions and the unfractionated soil (Total) from non-amended and amended soil (cross within symbols). Square: 0-50 μ m; Diamond: 50-200 μ m; Triangle: 200-2000 μ m; Inverted triangle > 2000 μ m. Standard deviations of each soil fraction correspond to 3 replicates of sampling dates.

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Fig. 5. Scores (A) and loadings (B) of the 2 main components of the PCA, representing 68% of the variability. Variables are relative abundance of C_{Straw} in the 12 FAME detected between Day 0 (white), Day 2 (grey) and Day 21 (black) of the incubation in each soil fractions and unfractionated soil (Total) from amended soil. Square: 0-50 µm; Diamond: 50-200 µm; Triangle: 200-2000 µm; Inverted triangle > 2000 µm. Standard deviations of each soil fraction correspond to 3 replicates of sampling dates.



Fig. 6. Evenness index of total FAME from each soil fraction, in control soil (grey bars) and
amended soil (dark grey bars) through time. Error bars represent the standard deviation of 3
experimental replicates.



Fig. 7. Evenness index of ¹³C-labelled FAME from each soil fraction of the amended soil through time. Error bars represent the standard deviation of 3 experimental replicates.







labelled C among soil fractions, after 2 and 21 days of incubation. Standard deviations of each 763

764 soil fraction correspond to 3 experimental replicates.